

### Lecture 3: Instabilities in the Solar Interior and in the Solar Atmosphere

#### 3.1. Review of the convectational instability in a hydrodynamic equilibrium medium

Let us consider a hydrodynamic equilibrium background medium, with background gas pressure  $p_0(z)$ , gas temperature  $T_0(z)$ , average mass density  $\rho_0(z)$ , and a relative uniform gravitational field  $g(-\hat{z})$ . The hydrodynamic equilibrium state yields

$$-\frac{dp_0(z)}{dz} - \rho_0(z)g = 0 \quad (3.1)$$

For ideal gas, we have

$$p_0(z) = \rho_0(z)RT_0(z) \quad (3.2)$$

Equation (3.2) yields,

$$\frac{1}{p_0} \frac{dp_0}{dz} = \frac{1}{\rho_0} \frac{d\rho_0}{dz} + \frac{1}{T_0} \frac{dT_0}{dz} \quad (3.2a)$$

Let us consider an air parcel, which moves from  $z(t=0) = z_0$  to  $z(t) = z_0 + \Delta z(t)$ , with thermal pressure  $p(t)$ , mass density  $\rho(t)$ , temperature  $T(t)$ , and initial conditions  $p(t=0) = p_0(z_0)$ ,  $\rho(t=0) = \rho_0(z_0)$ , and  $T(t=0) = T_0(z_0)$ . The horizontal pressure balance yields the gas pressure of the air parcel  $p(t)$  changes from  $p_0(z_0)$  to  $p_0(z_0 + \Delta z)$ . If this process is fast enough, then the air parcel will follow the adiabatic equation of state. i.e.,

$$\frac{d}{dt} \ln(p\rho^{-\gamma}) = 0 \quad (3.3)$$

where  $\gamma$  is the specific heat. The specific heat  $\gamma$  satisfies  $\gamma = (f + 2)/2$ , where  $f$  is the degree of freedom of the gas particles. For the ideal gas, we have  $p = \rho RT$ . Thus, the adiabatic equation of state can also be written as

$$\frac{d}{dt} \ln(p^{1-\gamma} T^\gamma) = 0 \quad (3.4)$$

Namely, the changes of the gas density and gas temperature satisfy the adiabatic equation of state. i.e.,

$$\frac{1}{p} \frac{dp}{dt} = \frac{\gamma}{\rho} \frac{d\rho}{dt} \quad (3.3a)$$

$$\frac{(1-\gamma)}{p} \frac{dp}{dt} + \frac{\gamma}{T} \frac{dT}{dt} = 0 \quad (3.4a)$$

The equation of motion of the air parcel when it moves to  $z = z_0 + \Delta z(t)$  is

$$\rho(t) \frac{d^2 z(t)}{dt^2} = \rho(t) \frac{d^2 \Delta z(t)}{dt^2} = - \left. \frac{dp_0}{dz} \right|_{z=z_0+\Delta z} - \rho(t)g \quad (3.5)$$

Equation (3.1) yields

$$\left. \frac{dp_0}{dz} \right|_{z=z_0+\Delta z} = -\rho_0(z_0 + \Delta z)g \quad (3.1a)$$

Substituting equation (3.1a) into equation (3.5), it yields

$$\frac{d^2\Delta z(t)}{dt^2} = \left( \frac{\rho_0(z_0 + \Delta z)}{\rho_0(t)} - 1 \right)g \quad (3.6)$$

We can conclude from equation (3.6) that, if  $\rho(t) > \rho_0(z_0 + \Delta z)$ , there will be a restoring force to move the air parcel back to  $z_0$ , and the system is stable to the convective instability. Or, if  $\rho(t) < \rho_0(z_0 + \Delta z)$ , the system would be unstable to the convective instability.

The numerator  $\rho_0(z_0 + \Delta z)$  in equation (3.6) can be written as

$$\begin{aligned} \rho_0(z_0 + \Delta z) &= \rho_0(z_0) + \Delta z \left. \frac{d\rho_0(z)}{dz} \right|_{z=z_0} + O(\Delta z^2 \frac{d^2\rho_0}{dz^2}) \\ &= \rho_0(z_0) \left\{ 1 + \Delta z \left. \frac{1}{\rho_0} \frac{d\rho_0}{dz} \right|_{z=z_0} + O(\Delta z^2 \frac{1}{\rho_0} \frac{d^2\rho_0}{dz^2}) \right\} \\ &= \rho_0(z_0) \left\{ 1 + \Delta z \left[ \frac{1}{\rho_0} \frac{dp_0}{dz} - \frac{1}{T_0} \frac{dT_0}{dz} \right]_{z=z_0} + O(\Delta z^2 \frac{1}{\rho_0} \frac{d^2\rho_0}{dz^2}) \right\} \\ &= \rho_0(z_0) \left\{ 1 + \Delta z \left[ \frac{-g}{RT_0} - \frac{1}{T_0} \frac{dT_0}{dz} \right]_{z=z_0} + O(\Delta z^2 \frac{1}{\rho_0} \frac{d^2\rho_0}{dz^2}) \right\} \end{aligned} \quad (3.7)$$

where equation (3.1) has been used to eliminate  $p_0(z_0)$  in equation (3.7).

Making use of equation (3.3), we can eliminate the denominator  $\rho(t)$  in equation (3.6).

Equation (3.3) yields  $p(t)\rho(t)^{-\gamma} = (p\rho^{-\gamma})_{t=0}$ , or

$$\begin{aligned} \rho(t) &= \left[ \frac{p(t)}{(p\rho^{-\gamma})_{t=0}} \right]^{1/\gamma} \\ &= \rho(t=0) \left[ \frac{p_0(z_0 + \Delta z)}{p_0(z_0)} \right]^{1/\gamma} \\ &= \rho_0(z_0) \left[ \frac{p_0(z_0) + \Delta z \left. \frac{dp_0}{dz} \right|_{z_0} + O(\Delta z^2 \frac{d^2p_0}{dz^2})}{p_0(z_0)} \right]^{1/\gamma} \\ &= \rho_0(z_0) \left[ 1 - \Delta z \frac{g}{RT_0(z_0)} + O(\Delta z^2 \frac{1}{p_0} \frac{d^2p_0}{dz^2}) \right]^{1/\gamma} \end{aligned} \quad (3.8)$$

where equation (3.1) has been used to eliminate  $p_0(z_0)$  in equation (3.8). Substituting equations (3.7) and (3.8) into equation (3.6), and ignoring the second order and the higher order terms, it yields

$$\begin{aligned}
 \frac{d^2 \Delta z}{dt^2} &\approx g \left\{ \frac{\rho_0(z_0) \left[ 1 + \Delta z \left( \frac{-g}{RT_0} - \frac{1}{T_0} \frac{dT_0}{dz} \right) \Big|_{z=z_0} \right]}{\rho_0(z_0) \left[ 1 - \Delta z \frac{g}{RT_0(z_0)} \right]^{1/\gamma}} - 1 \right\} \\
 &\approx g \left\{ \left[ 1 + \Delta z \left( \frac{-g}{RT_0} - \frac{1}{T_0} \frac{dT_0}{dz} \right) \Big|_{z=z_0} \right] \left[ 1 + \frac{1}{\gamma} \Delta z \frac{g}{RT_0(z_0)} \right] - 1 \right\} \\
 &\approx g \Delta z \left\{ \left( \frac{1}{\gamma} - 1 \right) \frac{g}{RT_0(z_0)} - \frac{1}{T_0} \frac{dT_0}{dz} \Big|_{z=z_0} \right\}
 \end{aligned} \tag{3.9}$$

If the background medium has an adiabatic-vertical-temperature profile  $T_{ad}(z)$ , i.e.,

$$p_0(z)^{1-\gamma} T_{ad}(z)^\gamma = \text{constant}$$

then

$$\left( \frac{1}{\gamma} - 1 \right) \frac{g}{RT_0(z_0)} = \frac{1-\gamma}{\gamma} \frac{1}{p_0} \frac{dp_0}{dz} = \left( \frac{1}{T_{ad}} \frac{dT_{ad}}{dz} \right) \tag{3.10}$$

Thus, we can define an adiabatic-vertical-temperature-change rate

$$\left( \frac{1}{T} \frac{dT}{dz} \right)_{ad} \equiv \left( \frac{1}{\gamma} - 1 \right) \frac{g}{RT_0(z_0)} \tag{3.11}$$

Substituting equation (3.11) into equation (3.9) it yields

$$\frac{d^2 \Delta z}{dt^2} \approx \Delta z \left\{ \left( \frac{1}{T} \frac{dT}{dz} \right)_{ad} - \frac{1}{T_0} \frac{dT_0}{dz} \Big|_{z=z_0} \right\} g = -\Delta z N^2 \tag{3.12}$$

where we have defined

$$N^2 \equiv \left\{ \left( -\frac{1}{T} \frac{dT}{dz} \right)_{ad} - \left( -\frac{1}{T_0} \frac{dT_0}{dz} \right) \Big|_{z=z_0} \right\} g \tag{3.13}$$

If  $N^2 > 0$ , the vertical displacement of the air parcel can lead to a vertical oscillation of the air parcel with angular frequency  $N$ , which is also called the Brunt-Väisälä frequency.

Namely, when the temperature-decreasing rate of the background gas  $-\frac{1}{T_0} \frac{dT_0}{dz} \Big|_{z=z_0}$  is less

than the adiabatic-temperature-decreasing rate  $\left( -\frac{1}{T} \frac{dT}{dz} \right)_{ad}$  the system is stable to the

convectonal instability. The system is unstable to the convectonal instability, if  $N^2 < 0$  or

if the temperature-decreasing rate of the background gas  $-\frac{1}{T_0} \frac{dT_0}{dz} \Big|_{z=z_0}$  is greater than the

adiabatic-temperature-decreasing rate  $\left( -\frac{1}{T} \frac{dT}{dz} \right)_{ad}$ .

### 3.2. The convectonal instability in the solar interior

If we consider the interior of the sun as a spherically symmetric hydrodynamic equilibrium background medium, with background plasma pressure  $p_0(r)$ , plasma temperature  $T_0(r)$ , plasma number density  $n_0(r)$ , average ion mass  $\bar{\mu}(r)$ . The average mass density is

$$\rho_0(r) = n_0(r)\bar{\mu}(r) \quad (3.14)$$

The gravitational field  $\mathbf{g}(r) = -\hat{r}g(r)$  can be determined by

$$-\hat{r}g(r) = -\hat{r}4\pi \int_0^r r'^2 \rho_0(r') dr' \quad (3.15)$$

The hydrodynamic equilibrium state yields

$$-\frac{dp_0(r)}{dr} - \rho_0(r)g(r) = 0 \quad (3.16)$$

For ideal gas, we have

$$p_0(r) = n_0(r)k_B T_0(r) = \rho_0(r) \frac{k_B}{\bar{\mu}(r)} T_0(r) \quad (3.17)$$

Based on solar seismology, scientists can determine the radial distribution of the acoustic wave speed  $C_{s0}(r)$  in solar interior, where

$$C_{s0}(r) = \sqrt{\frac{\gamma p_0(r)}{\rho_0(r)}} = \sqrt{\gamma k_B} \sqrt{\frac{T_0(r)}{\bar{\mu}(r)}} \quad (3.18)$$

The standard solar model (xxxx, 199x), provide an theoretical model for the average ion mass distribution in the solar interior  $\bar{\mu}(r)$ . Thus, the five equations (3.14), (3.15), (3.16), (3.17) and (3.18) together with a given  $\bar{\mu}(r)$  model and observed  $C_{s0}(r)$ , we can determine the five unknowns  $p_0(r)$ ,  $T_0(r)$ ,  $n_0(r)$ ,  $\rho_0(r)$ , and  $g(r)$ .

If we include the magnetic force in the convection zone, the equilibrium state along the radial direction becomes

$$-\frac{\partial p_0}{\partial r} - \rho_0(r)g(r) + (\mathbf{J} \times \mathbf{B})_r \approx -\frac{\partial p_0}{\partial r} - \rho_0(r)g(r) - \frac{\partial}{\partial r} \frac{B^2}{2\mu_0} + \frac{\mathbf{B} \cdot \nabla}{\mu_0} B_r = 0 \quad (3.16a)$$

Since the magnetic field cannot be spherical symmetry, if the magnetic force is included as described in equation (3.16a) the average ion mass  $\bar{\mu}$  should become a spherical asymmetric function, i.e.,  $\bar{\mu} = \bar{\mu}(r, \theta)$ . Namely, the spherical symmetric model of  $\bar{\mu}(r)$  is no longer applicable to the plasma in the solar convection zone. Likewise, by including the magnetic force, it might lead to  $p_0(r, \theta)$ ,  $T_0(r, \theta)$ ,  $n_0(r, \theta)$ , in the solar convection zone. Thus, the convectonal instability in the solar convection zone is far more complex than the convectonal instability in the Earth's troposphere.